

Name \_\_\_\_\_

Date \_\_\_\_\_

Partners \_\_\_\_\_

Section \_\_\_\_\_

## Atmospheric Stability

**Purpose:**

Use vertical soundings to determine the stability of the atmosphere.

**Equipment:**

Station Thermometer

Min./Max. Thermometer

Anemometer

Barometer

Psychrometer

Rain Gauge

Psychometric Tables

Barometric Correction Tables

**I. Surface observation.**

Begin the first 1/2 hour of lab performing a surface observation. Make sure you include pressure (station, sea level, and altimeter setting), temperature, dew point temperature, wind (direction, speed, and characteristics), precipitation, and sky conditions (cloud cover, cloud height, & visibility). From your observation generate a METAR and a station model.

A. Generate a METAR for today's observation

B. Generate a station model for today's observation.

**II. Plot the sounding**

On the skew T-ln p plot the CHS sounding. Locate the lifting condensation level (LCL), the level of free convection (LFC), and the equilibrium level (EL).

1. The LCL is found by following the dry adiabat from the surface temperature and the saturation mixing ratio from the dew point temperature. Where these two paths intersect, we would have condensation taking place. This is the LCL and represents the base of the cloud.
2. The LFC is found by following a moist adiabat from the LCL until it intersects the temperature sounding. Beyond this level a surface parcel will become warmer than the environment and will freely rise.
3. The EL is found by following a moist adiabat from the LFC until it intersects the temperature sounding again. At this point the surface parcel will be cooler than the environment and will no longer rise on its own accord.

### III. Stability of individual layers of a sounding

The stability of the atmosphere depends on the environmental lapse rate and the humidity in the air. The lapse rate is categorized as either absolutely stable, saturated neutral, conditionally stable, dry neutral, and absolutely unstable. Our goal in this exercise is to analyze the stability of each reported layer in the vertical sounding. The sounding data is given for CHS (Charleston, SC) 0000Z May 1, 1998.

1. **Calculate the lapse rate** - Calculate the temperature difference between adjacent reported levels of the sounding and divide by the height difference.

$$\gamma = -\frac{T_2 - T_1}{h_2 - h_1}$$

After calculating the lapse rate record the results on the data sheet.

2. **Assign stability to each level of the sounding** - Look at the lapse rate between each reported level of the sounding as calculated in part 1. Compare this with the plotted data on the included skew T-ln p. The plot will not correspond exactly to the layers over which we calculated because only the mandatory levels have been maintained. Look at the plotted sounding and determine if this lapse rate is less than, equal to, or greater than the moist adiabatic lapse rate,  $\gamma_m$ , at that level. Do the same comparison with the dry adiabatic lapse rate,  $\gamma_d$ , at that level. Assign a stability based on the following criteria.

$\gamma > \gamma_d$ and $\gamma > \gamma_m$	au (absolutely unstable)
$\gamma = \gamma_d$	d (dry neutral)
$\gamma < \gamma_d$ and $\gamma > \gamma_m$	cu (conditionally unstable)
$\gamma = \gamma_m$	m (moist neutral)
$\gamma < \gamma_d$ and $\gamma < \gamma_m$	as (absolutely stable)

(In class we stated that the dry adiabatic lapse rate is 10 °C and the moist adiabatic lapse rate is 6 °C. These are estimates near the surface. To get better results use the skew T-ln p included with this lab.)

3. **Evaluate overall stability of each level of the sounding** - Look at the different levels and determine if they are saturated or dry. Under the following conditions the air will be considered saturated.

$T - T_d \leq 2 \text{ }^\circ\text{C}$	when $T > 0 \text{ }^\circ\text{C}$
$T - T_d \leq 3 \text{ }^\circ\text{C}$	when $-10 \text{ }^\circ\text{C} < T < 0 \text{ }^\circ\text{C}$
$T - T_d \leq 4 \text{ }^\circ\text{C}$	when $T \leq -10 \text{ }^\circ\text{C}$

If the level is conditionally unstable and it is saturated, then the level is unstable. Go through all of the levels and declare them either stable or unstable.

DATA SHEET  
Stability of layers

CHS 0000Z May 1, 1998

P <sub>1</sub> (mb)	P <sub>2</sub> (mb)	γ (°C/km)	Stability	Saturated	Overall Stability
1011	1009				
1009	1000				
1000	975				
975	925				
925	862				
862	850				
850	778				
778	700				
700	652				
652	604				
604	517				
517	508				
508	500				

Date:0000Z 1 MAY 98  
Station: CHS  
WMO ident: 72208  
Latitude: 32.90  
Longitude: -80.03  
Elevation: 14.00

LEV	PRES mb	HGHT m	TEMP C	DEWP C	RH %	DD C	WETB C	DIR deg	SPD knt	THETA K	THE-V K	THE-W K	THE-E K	W g/kg
SFC	1011	14	20.0	18.3	90	1.7	18.9	90	11	292.2	294.6	291.6	329.6	13.20
1	1009	32	20.2	16.9	81	3.3	18.0	90	11	292.6	294.8	290.8	326.9	12.09
2	1000	112	19.4	17.6	89	1.8	18.2	90	12	292.5	294.8	291.3	328.8	12.76
3	975	330	17.2	16.7	97	0.5	16.9	94	13	292.5	294.7	291.0	327.5	12.35
4	925	779	14.8	14.3	97	0.5	14.5	80	6	294.4	296.4	290.7	326.4	11.14
5	862	1375	12.6	12.1	97	0.5	12.3	256	11	298.1	300.0	291.2	328.3	10.34
6	850	1493	12.4	9.7	84	2.7	10.7	260	12	299.1	300.8	290.3	325.3	8.92
7	778	2232	9.2	3.2	66	6.0	5.9	260	15	303.4	304.5	289.4	322.0	6.20
8	700	3100	3.8	2.6	92	1.2	3.1	250	21	306.7	307.9	290.8	326.8	6.61
9	652	3673	0.6	-5.4	64	6.0	-2.1	250	26	309.4	310.1	289.3	321.7	3.92
10	604	4284	-1.5	-10.5	50	9.0	-5.2	250	30	313.8	314.3	289.6	323.0	2.84
11	517	5502	-10.1	-20.1	44	10.0	-13.1	215	32	317.7	318.0	289.6	322.7	1.50
12	508	5638	-9.7	-34.7	11	25.0	-14.9	215	33	319.8	319.8	289.1	321.2	0.40
13	500	5760	-10.1	-41.1	6	31.0	-15.6	215	34	320.7	320.8	289.2	321.5	0.21